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1 A novel framework to study the effect of tree architectural traits on stemflow yield and its
2 consequences for soil-water dynamics

3
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10
11 Abstract

12 A novel experimental approach and numerical framework are proposed to study the effect of tree
13 architectural traits on stemflow yield and its effects on soil-water dynamics. The framework includes
14 a data mining workflow employing information from two experimental steps: (i) evaluation of the
15 effect of tree aboveground architecture on stemflow yield and (ii) quantification of specific
16 parameters for soil-water dynamics with and without stemflow. We studied double-funnelling
17 (stemflow and root-induced preferential flow) under three sycamore (*Acer pseudoplatanus* L.) trees
18 growing on a slope in Scotland during the summer season and measured architectural traits. Stemflow
19 yield ranged from 1.3 to 3.8 % of the incident rainfall, with funnelling ratios of between 2.2 ± 2.1 and
20 5.2 ± 3.9 . Double-funnelling to a depth of up to 400 mm beneath the soil surface occurred as matrix
21 flow and was significantly and positively correlated with the vertical root distribution. Soil-water
22 dynamics were distinctly different with and without stemflow. Our framework revealed that the
23 number of tree branches, their insertion angle, leaf number, and stem basal diameter influenced
24 stemflow yield within rainfall thresholds of 1.1 and 3.5 mm d⁻¹. The framework also showed that
25 stemflow yield had a negative impact on soil matric suction, while air temperature was the most
26 influential covariate affecting soil-water dynamics, likely due to its strong correlation to
27 evapotranspiration during the summer season. In spite of the study limitations, such as small sample
28 size and differences between individuals, we show that the proposed framework and experimental
29 approach can contribute to our knowledge of how stemflow generated aboveground triggers major
30 responses in soil-water dynamics belowground.

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34 Keywords: temperature, soil moisture, matric suction, preferential flow, sycamore

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37 1. Introduction

38

39 Stemflow is a poorly studied hydrological process that occurs mostly in woodland
40 ecosystems (Levia and Germer, 2015). Stemflow comprises the funnelling of incident
41 precipitation around the tree stem, with subsequent flow occurring along roots and into soil
42 pores (i.e. double-funnelling; Johnson and Lehmann, 2006). Stemflow has been overlooked
43 due to its point-source nature and its apparently small contribution to the water cycle (Levia
44 et al., 2011). However, many geoscientists now acknowledge the potential role of stemflow
45 in the regulation of hydrological and biogeochemical cycles in forests and shrublands (Levia
46 and Germer, 2015).

47

48 Stemflow generally represents 1 to 20 % of incident precipitation (Levia and Frost, 2003).
49 But, in reality, a substantial volume of water can concentrate around the tree bole in a single
50 precipitation event (e.g. Gonzalez-Ollauri and Mickovski, 2017a). Stemflow is formed in the
51 forest canopy, and recent work has strived to unveil how canopy architecture can regulate the
52 formation of stemflow before being funnelled belowground (e.g. Levia et al., 2010, 2015;
53 Bialkowski and Buttle, 2015; Yuan et al., 2017; del Campo et al., 2018). However, tree
54 architecture is complex (e.g. Barthélémy and Caraglio, 2007; Côté et al. 2011), and varies
55 with tree age (e.g. Meinzer et al., 2011), and across species and biomes (e.g. Wright et al.,
56 2005; Kattge et al., 2011). Outcomes from advanced data mining frameworks (e.g. Torgo,
57 2011) could have important applications in the elucidation of strategies by practitioners
58 seeking to control relevant phenomena that are regulated through stemflow (e.g. groundwater
59 recharge and nutrient cycling, Levia et al., 2011; del Campo et al., 2014; 2018).

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61 Double-funnelling results into various modes of subsurface flow, e.g. preferential flow,
62 bypass flow and matrix flow (Liang et al., 2011; Spencer and van Meerveld, 2016), or even
63 surface runoff (Herwitz, 1986; Miyata et al., 2009). Tracking double-funnelling is labour-
64 intensive (e.g. Liang et al., 2011; Germer, 2013; Spencer and van Meerveld, 2016), and it is
65 difficult to identify the subsurface flow mode, its potential drivers, and the subsequent effect
66 in the soil. For trees growing on slopes, stemflow is funnelled towards the downslope part of
67 the stem and soil (Liang et al., 2011; Spencer and van Meerveld, 2016), where macropores
68 and gaps between the root and soil channel this subsurface flow (Martinez-Meza and
69 Whitford, 1996; Schwärzel et al., 2012; Spencer and van Meerveld, 2016). There is little
70 convincing information, however, indicating whether the subsurface flow mode associated to
71 double-funnelling is species-specific and season-dependent, or whether it changes with

meteorological conditions, e.g. rainfall intensity or inter-rainfall interval (van Stan et al., 2012; Tanaka et al., 2017). The subsurface flow mode may also depend on soil characteristics, such as moisture content and texture, as well as plant root density. In spite of the advances made over the last decade to generate knowledge on how soil hydrology is impacted by stemflow (for review, see Levia and Germer, 2015), linking double-funnelling and subsequent changes to soil-water dynamics in response to meteorological events has not yet been attempted.

Changes in soil moisture and matric suction related to double-funnelling can help to clarify what level of stemflow yield leads to noticeable hydrological responses in soil. Measuring soil hydrological variables also allows an assessment of whether double-funnelling occurs along the same soil paths repeatedly, and whether it influences mechanical stresses in soil, that could then affect how vegetation protects against rainfall-induced, shallow landslides (Gonzalez-Ollauri and Mickovski, 2017a,b). Changes in soil temperature at zones with signs of double-funnelling may also indicate the arrival of water and nutrients transported from the canopy to soil (e.g. Germer et al., 2012), that in turn alter soil microbial activity (McClain et al., 2003; Kuryakov and Blagodatskaya, 2015; Rosier et al., 2016). However, studying soil-water dynamics influenced by stemflow is experimentally challenging and data are difficult to analyse. Therefore, robust experimental and numerical frameworks should enable us to overcome these problems and so elucidate relationships between stemflow and soil-water dynamics.

The aim of our study is to propose a novel experimental approach and numerical framework to estimate how aboveground tree architecture affects stemflow yield and, subsequently, soil-water dynamics. Our objectives are: (i) evaluation of the effect of aboveground tree architecture on stemflow yield; (ii) identification of soil-root zones subjected to double-funnelling; (iii) quantification of soil moisture, matric suction and soil temperature with and without stemflow and (iv) evaluation of the effect of two external meteorological variables (i.e. rainfall and air temperature) on soil-water dynamics in contrast with stemflow yield. To achieve these objectives, we studied double-funnelling and its effect on soil-water dynamics under sycamore (*Acer pseudoplatanus* L.) trees. Sycamore is a deciduous, broadleaf, fast-growing tree with a broad, domed crown and smooth bark (Pasta et al., 2016). Mature sycamores are very resistant to wind loading, coastal exposure, and pollution, and these features make it a useful species for protecting slopes against landslides, erosion, and rockfall

(Norris et al. 2008). Therefore, sycamore is a potentially useful species in protection forests and on unstable slopes. Understanding how the architectural traits of sycamore affect soil hydrological characteristics will allow a full assessment of the utility of this species for ecological engineers.

2. Materials and Methods

2.1. Tree individuals and study site

The study site was located adjacent to Catterline Bay, Aberdeenshire, UK (WGS84 Long: -2.21 Lat: 56.90; supplementary material Fig. S1a), within the temperate humid climate zone (Cgc: subpolar oceanic climate; Köppen, 1884). The mean annual temperature over the period 2011 – 2018 was 8.9 °C and the mean annual rainfall was 556.8 mm (voor de Poorte, 2018). The precipitation at the site is characterised by frequent, low-intensity rainfall events (Gonzalez-Ollauri and Mickovski, 2016; 2017a) and the prevailing wind is south-westerly. Three adjacent sycamore (*Acer pseudoplatanus* L.) individuals (Syc1, Syc2, and Syc3), that were approximately 40 years old, were chosen for the study (Fig. S1b). The stand was established on a $20.3 \pm 11.6^\circ$ southwest facing slope, oriented $260\text{--}280^\circ$ from due north. Soil is well-drained (saturated hydraulic conductivity: $7.1 \times 10^{-5} \text{ m s}^{-1}$), with a soil organic matter content of 12.4 %. Landslide prone silty sands (sand: 66.8 %; silt: 1.4 %; clay: 0.8 %) overlie conglomerate rock. The topsoil at the site (to a depth of 600 mm beneath the soil surface) had a mean dry bulk density of 1.3 g cm^{-3} , a drained cohesion of 30.5 kPa, a mean angle of internal friction of 19.4° , and a specific gravity of 2.6.

2.2. Aboveground tree architectural traits

The aboveground architecture of the tree individuals was characterised with 13 readily measurable traits reported to influence stemflow yield (Levia et al., 2015; Fig. 1). A surveyors' meter tape was used to measure: (i) the tree's basal area or diameter at breast height (*DBH*; m) and (ii) the projected canopy-crown area (C_A ; m^2) according to Spoke's distance method (Blozan, 2008), assuming a circular crown projection on the ground surface. A spherical crown densiometer was used to estimate (iii) the canopy cover fraction (c ; %; Lemon, 1956). (iv) Tree height (H_t ; m) was measured from an upslope position with a Leica®

laser meter. (v) The leaf area index (LAI) was estimated with the Wolf et al. (1972) direct method. A hand-held inclinometer was used to measure (vi) the stem lean from the vertical (S_L ; °), (vii) the maximum branch insertion angle from the horizontal ($_{mx}Br_a$; °), and (viii) the mean branch insertion angle above the horizontal ($_{av}Br_a$; °). (ix) The number of primary (PBr ; developing directly from the main stem) and (x) secondary branches (SBr ; borne directly on the primary branches) were counted manually to then calculate the branch count per unit projected area of tree crown (Levia et al., 2015). Finally, (xi) the total number of leaves (nL ; m^{-2}), (xii) leaf biomass (L_{BM} ; $g\ m^{-2}$), and (xiii) branch biomass (Br_{BM} ; $Kg\ m^{-2}$) per unit projected area of tree crown was estimated through a destructive method that involved cutting a primary and two secondary branches from a fourth, unstudied, sycamore individual. Leaves were counted and, along with woody parts, were oven-dried at 60°C until a constant mass was reached.

Figure 1. Illustration showing tree architectural traits measured in this study. DBH : diameter at breast height; C_A : projected canopy-crown area; c : canopy cover fraction; Ht : tree height; LAI : leaf area index; S_L : stem lean; Br_a : branch insertion angle. On the right-hand side, the concept of LAI and two c examples are illustrated, i.e. dense canopy crown with $c=100\%$ and sparse canopy crown with $c=50\%$, where the white squares portray the penetration of sunlight through the canopy.

2.3. Double-funnelling and subsurface flow

To identify the zones where double-funnelling results in an accumulation of subsurface flow, on 18/06/2018, we sprayed 20 l of a dye solution (i.e. Brilliant Blue FCF; $5\ g\ l^{-1}$) between ground level and a height of 1.7 m on the downslope side of the stem (Laing et al., 2011; Nespoulous et al., 2019) on two of the sycamore individuals (Syc1 and Syc2; Fig. 2b). For this, we used a 20 l backpack sprayer for 35 minutes per individual, corresponding to a

precipitation intensity of 45.7 mm h^{-1} . Before we sprayed the dye solution, we used a moisture profile probe (Delta-T®) to measure the mean soil volumetric moisture content (θ_v , %) within a soil depth range of 0.0 – 0.3 m below ground level (b.g.l) and 0.15 m away from the tree bole on the downslope side of the stem. Mean profile θ_v was $5.1 \pm 2.3\%$. Thirty-six hours after spraying the dye solution, we dug a 2.0 x 1.0 m trench downslope from each tree 0.3 m away from the tree bole (Fig. 2). The wall of each trench was smoothed with a knife prior to mapping the dyed areas and the root profiles onto a 1.0 m x 0.5 m gridded acetate sheet (Böhm, 1979). The area of dyed soil was quantified at 0.1 m intervals along the soil profile by examining the proportion inside each grid square that was stained (Nespoulous et al., 2019). The cross-sectional area of all roots (A_r ; mm^2) was quantified at the same soil depth intervals (Gonzalez-Ollauri and Mickovski (2016); Eqs. 1 and 2 (Table 1)), once sycamore roots had been identified (Reward et al., 2012).

Figure 2. (a) Illustration showing the experimental setup deployed onsite to study soil-water dynamics with stemflow (Syc2) and without stemflow (Syc1); (b) Brilliant Blue FCF was sprayed on the stem of Syc1 and Syc2 prior to suppressing stemflow in Syc1 with a gutter. Most of the dye solution infiltrated the soil next to the tree bole; (c-e) trenches were dug to observe the distribution of Brilliant Blue FCF belowground and to install sensors monitoring soil-water dynamics at the root-soil interface. The trenches were covered with a landscape mat to avoid interference with dripfall and throughfall, θ_v-1 , θ_v-2 , and θ_v-3 indicate the soil moisture probes used under Syc1 (see Table 2). More images from the experimental set up are shown in supplementary material.

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Definition	Equation	No	Parameters	Units	Source
Cross-sectional area of roots	$Ar = \pi \cdot (\Sigma d/2)^2$	Eq.(1)	Ar : Cross-sectional area of all roots at a given soil depth d : root diameter	mm ² mm	Gonzalez-Ollauri and Mickovski (2016)
Vertical root distribution	$Ar(z) = Aro \cdot e^{-z/b}$	Eq.(2)	$Ar(z)$: cross-sectional area of all roots along the soil profile Aro : cross-sectional area of the stump b : mean rooting depth z : soil depth	mm ² mm ² mm mm	Preti et al. (2010) Gonzalez-Ollauri and Mickovski (2016)
Stemflow yield	$Sy = a + b \cdot Pg$	Eq.(3)	Sy : stemflow yield Pg : incident rainfall a : intercept b : slope	L mm	Deguchi et al. (2006)
Stemflow funnelling ratio	$SFR = \frac{Sy}{Pg \cdot DBH}$	Eq.(4)	DBH : stem diameter at breast height (i.e. 1.4 m from the ground level)	m ²	Herwitz (1986)
[‡] Stemflow yield and tree architecture	$Sy = f(Pg, Ht, C_A, DBH, c, LAI, S_L, PBr, SBr, mxBra, mnBra, nL, L_{BM}, Br_{BM}, C_{BM})$	Eq.(5)	Ht : tree height C_A : projected canopy-crown area DBH : diameter at breast height (i.e. 1.4 m from the ground level) c : canopy cover fraction LAI : leaf area index S_L : stem lean from the vertical axis measured at the ground level [†] PBr : number of primary branches (i.e. developing from the main stem; Fig. 1a) [†] SBr : number of secondary branches (i.e. borne on the primary branches; Fig. 1a) $mxBra$: maximum branching angle from the horizontal axis $mnBra$: mean branch angle from the horizontal axis [†] nL : leaf count [†] L_{BM} : leaf biomass [†] Br_{BM} : total branch biomass (i.e. primary and secondary branch biomass) [†] C_{BM} : crown biomass	m m ² m % m ² m ⁻² ° m ⁻² m ⁻² ° ° m ⁻² g m ⁻² g m ⁻² g m ⁻²	This study
[‡] Soil temperature	$ST = f(Pg, Sy, Ta)$	Eq.(6)	ST : soil temperature measured from June to October, 2018 Pg : incident rainfall (i.e. rainfall that reaches the soil without vegetation) measured from June to October, 2018 Sy : stemflow yield measured from June to October, 2018 Ta : air temperature measured from June to October, 2018	° mm mL °	This study
[‡] Soil moisture	$\theta v = f(Pg, Sy, Ta)$	Eq.(7)	θ_v : soil moisture measured from June to October, 2018	/1	This study

Soil matric suction	$\psi = f(Pg, Sy, Ta)$	Eq.(8)	Ψ : matric suction measured from July to August, 2018	kPa	This study
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2.4. Quantification of stemflow and funnelling ratio

Stemflow was suppressed for two of the three sycamore individuals (Syc1 and Syc3, Figs. 2a,b), so that the influence of aboveground tree architecture on double-funnelling (Section 2.6; Fig. 3a) could be quantified. Suppressing stemflow also allowed us to determine indirectly its contribution to soil-water dynamics (Fig. 3a). This suppression was achieved by sealing one stemflow gutter (32 mm diameter, corrugated vinyl tube, with a third of its perimeter cut off to enable the collection of stemflow), starting at a height of 1.7 m up the tree stem, and revolving one and a half times around the stem (Figs. 2a,b; Gonzalez-Ollauri and Mickovski, 2017a). Each stemflow gutter terminated in a 25 l opaque plastic container where stemflow water was collected and stored until measurement (Fig. 2a). The amount of stemflow collected was measured with a graduated cylinder on a rainfall event basis from July to October, 2018. On 21/06/2018, and after trenches had been dug (Section 2.3), we cleared out the understory vegetation, and covered the ground surface below the canopies of Syc1 and Syc2 with a landscape mat. This mat prevented the infiltration of water from dripfall and throughfall (e.g. Zimmermann and Zimmermann, 2014) into the soil but allowed stemflow infiltration under Syc2 (Fig. 2c).

Stemflow yield was compared against incident rainfall (P_g ; mm d⁻¹) by fitting linear regression models (Eq. 3, Table 1; Deguchi et al., 2006) in R v3.5.1 (R Core Team, 2018). Rainfall and air temperature time series were retrieved with a 1-minute resolution from a Davis Vantage Pro2 meteorological station located *in situ* and away from the canopy's influence (voor de Poorte, 2018; Fig. S1a). We monitored 35 rainfall events in total, with an event defined as having a minimum depth of 0.4 mm and being separated by at least 2 hours without rainfall. We strived to measure stemflow volume 2h after rainfall or the following morning when the former was not possible. When more than one rainfall event occurred before we could measure stemflow (seven events in total), we assumed a linear relationship between rainfall and stemflow, and we proceeded as follows: (i) we measured the total stemflow volume collected in the container; (ii) we then discretised rainfall events by aggregating the 1 minute rainfall time steps into hourly steps; (iii) we subset and pooled consecutive time steps with rainfall depths above 0.4 mm; (iv) we summed up the rainfall depth for the pool of unmonitored events; (v) we estimated the volume ratio for each event considering the total rainfall volume, and finally, (vi) we applied this ratio to the total

stemflow volume collected in the container. Once stemflow yield was known, the stemflow funnelling ratio (*SFR*) was calculated at the individual level as indicated in Eq. 4 (Table 1; Herwitz, 1986).

2.5. Influence of stemflow and its suppression on soil-water dynamics

After trenches had been dug for the observation of dyed water originating from stemflow (Section 2.3), we monitored soil-water dynamics under both stemflow (Syc2) and suppressed stemflow (Syc1). We also monitored soil-water dynamics in zones with and without signs of double-funnelling (i.e. with and without blue staining) (Figs. 2d,e). To do so, we measured soil temperature (*ST*, °C), soil volumetric moisture content (θ_v , / l), and soil matric suction (ψ , kPa) using automatic sensors during the growing season of 2018 (i.e. late June – early October, 2018).

Every sensor was deployed at the interface between roots larger than 5 mm in diameter and soil (Fig. 2a) at locations noted in Table 2. Soil moisture content was monitored with seven time-domain reflectometry sensors (TDR; CS616 – Campbell Scientific, UK) installed at different soil depths (Table 2) in the excavated trenches (Figs. 2c, d). Soil temperature was monitored in the excavated trenches under stemflow and suppressed stemflow by installing one temperature probe (T107 – Campbell Scientific, UK) per trench and in areas that had been stained with dye (Table 2; Figs. 2c,d). Soil matric suction was monitored with two field tensiometers/piezometers (T4 – UMS GmbH, Germany) installed vertically within the soil-root zone (i.e. 0-500 mm beneath the soil surface; Gonzalez-Ollauri and Mickovski, 2017a; Tardio et al., 2016) and, at 0.1 m from the downslope side of the tree boles of Syc1 and Syc2. All sensors were wired to a solar powered CR-1000 data logger (Campbell Scientific, UK), that collected records for *ST*, θ_v , and ψ at 15 minute time steps. To test the operational capacity of the instrumental setup, we undertook four stemflow simulation events at the onset of the monitoring period (event 1 on 28/06/2018; event 2 on 03/07/2018; event 3 on 04/07/2018; and event 4 on 09/07/2018). Each stemflow simulation event consisted of spraying 20 L of tap water at a height of 1.7 m over the downslope side of the stem of Syc1 and Syc2 using a backpack sprayer for 35 minutes (i.e. rainfall intensity of 45.7 mm h⁻¹). The results from these simulations were analysed together with the records derived from real stemflow events occurring during the monitoring period.

Table 2. Sensor type and location in stained/unstained soil under trees with stemflow or where stemflow was suppressed. Sensors measure soil temperature (ST), soil moisture (θ_v) and matric suction (ψ). Sensor Id refers to the sensor type and number.

Variable	Sensor	Sensor Id	Tree	Treatment	Positioned in zone previously stained with blue dye from stemflow?	Depth (mm)
Soil temperature	T107	$ST1$	Syc1	Suppressed stemflow	Yes	150
	T107	$ST2$	Syc2	Stemflow	Yes	150
Soil moisture	CS616	θ_v-1	Syc1	Suppressed stemflow	Yes	100
	CS616	θ_v-2	Syc1	Suppressed stemflow	No	100
	CS616	θ_v-3	Syc1	Suppressed stemflow	Yes	260
	CS616	θ_v-4	Syc2	Stemflow	No	400
	CS616	θ_v-5	Syc2	Stemflow	Yes	100
	CS616	θ_v-6	Syc2	Stemflow	No	100
	CS616	θ_v-7	Syc2	Stemflow	Yes	400
Soil matric suction	T4	$\psi-1$	Syc1	Suppressed stemflow	Yes	300-400
	T4	$\psi-2$	Syc2	Stemflow	Yes	300-400

2.6. Framework description and implementation

We defined a cascade-style, numerical framework that incorporates a data mining workflow to evaluate the relationships between tree architecture, stemflow, and stemflow-derived soil-water dynamics (i.e. soil temperature, volumetric soil moisture content and matric suction, Fig. 3a). The framework was built using the statistical software R v3.5.1 (R Core Team, 2018). Data mining was implemented twice (Figs. 3a-b): (i) to explore the effect of tree architecture on stemflow yield, and (ii) to assess the effect of stemflow yield, rainfall, and air temperature on soil-water dynamics.

The data mining workflow fits 100 boosted regression tree models (BRTs; Breiman et al. 1984) without pruning, using the R package “rpart” (Therneau and Atkinson, 2018). Each BRT was fitted using a training dataset generated through a bootstrapping method with

replacement (Efron, 1979). Accordingly, a dataset containing 70 % of the observations (i.e. the training dataset) was extracted for each model run. The model outcomes were then cross-validated with the remaining 30 % of the data. Model quality was evaluated depending on the value of the coefficient of determination (R^2) and the root mean square error (RMSE), between predicted and observed values, following the least-squares criterion (e.g. Bruce and Bruce, 2017). The best fitted BRT was selected on the basis of its quality and the amount of information carried by the regression tree (i.e. the number of relevant covariates portrayed in the BRT). In addition, the relative influence (RI ; %; e.g. del Campo et al., 2018) of each covariate on the response variable was examined through the evaluation of the decrease in node impurities (i.e. reduction in mean squared error), produced by splitting each regression tree on a given metric using the R package “caret” (Khun, 2018). The nature of the relationships between response and predictor variables were then evaluated on the basis of partial dependence plots (e.g. Tanaka et al., 2017), retrieved with the R package “pdp” (Greenwell, 2017).

The steps followed to implement the proposed framework (Fig. 3a) are illustrated in Fig. 3b. Firstly, the data mining workflow was implemented to fit stemflow yield BRT models using Eq.5 (Table 1) and the information collected from Syc1 and Syc3. Incident rainfall was included as covariate in the fitting of BRTs (Fig. 3a). The purpose of this was twofold: (i) to enable prediction of stemflow yield under varying incident rainfall, and (ii) to investigate the effect of tree architecture masked by the relationship between incident rainfall and stemflow yield (Deguchi et al., 2006) – i.e. how rainfall events are partitioned into stemflow by an array of steady-state, tree architectural traits at an individual level. Model quality, covariates’ RI , and relationship between predictors and response variables were then examined as described above (Fig. 3a). Following model cross-validation and best model selection (Fig. 3a), stemflow yield was predicted for Syc2 with the best performing BRT. For this, the measured tree architectural traits in Syc2 and incident rainfall were used as inputs (Fig. 3b).

Subsequently, soil-water BRT models were fitted to the monitored soil-water variables (i.e. ST , θ_v , and ψ), using Eqs. 6, 7 and 8, respectively (Table 1). To this end, air temperature, incident rainfall, and predicted stemflow yield were used as inputs for Syc2 (Fig. 3b), while the stemflow yield inputs were assumed to be zero for Syc1 (i.e. tree with suppressed stemflow). Air temperature was included to investigate soil-water dynamics during drying (i.e. in the absence of precipitation), as air temperature and soil-water tend to be tightly

coupled (e.g. Feng and Liu, 2015). Incident rainfall was included to detect soil-water dynamics that were not related to stemflow in the context of this study, e.g. subsurface flow generated away from the canopy's influence, but still reaching the soil-root zone under study (Bogaard, 2001). BRT models were fitted per monitoring sensor (Table 2) and per tree individual where soil-water dynamics were studied (Syc1 and Syc2).

All time series were aggregated into daily time steps prior to fitting the BRT models. Model quality was evaluated using the probability density functions from the pool of R^2 values retrieved from cross-validation (Fig. 3a). The *RI* of the covariates (i.e. air temperature, incident rainfall, and stemflow yield), on soil-water dynamics and the relationship between predictors and response variables were then examined as described above (Fig. 3a).

Figure 3. (a) Proposed framework to study the effect of tree architecture on stemflow yield, and the effect of stemflow yield and external meteorological variables (i.e. rainfall and air temperature) on soil-water dynamics, i.e. soil temperature, soil moisture, and matric suction. (b) Flow chart illustrating the implementation of the proposed framework in this study. See online version for colours.

2.7. Statistical analysis

The vertical distribution of tree roots and the soil area affected by double-funnelling were correlated using Pearson's correlation (r) tests. Non-parametric Kruskal-Wallis (χ^2) tests

were then performed to determine differences in stemflow yield between individuals on an event basis. Kruskal-Wallis tests were also used to evaluate differences in daily levels of ST , θ_v , and ψ between trees with stemflow and those where stemflow was suppressed. Differences in model quality between the fitted BRTs (i.e. R^2 ; Section 2.6) as well as differences between covariates' RI generated for the 100 BRTs fitted to tree architectural traits and to soil-water dynamics parameters, respectively, were also evaluated with Kruskal-Wallis tests. Stemflow yield was excluded from statistical analyses evaluating differences between covariate's RI on ST , θ_v , and ψ for Syc1, as stemflow yield was assumed to be zero for this tree (i.e. tree with suppressed stemflow). All statistical tests were performed using the software R v3.5.1 (R Core Team, 2018), at 95% and 99% confidence levels, following normality testing through Shapiro-Wilk tests.

3. Results

3.1. Aboveground tree architectural traits

Sycamore trees had smooth stems with 3 to 5 primary branches inserted above 1.7 m, from which 60 to 80 secondary branches emerged (Table 3). The two individuals on which stemflow yield was quantified (Syc1 and Syc3; Table 3) were similar with regard to DBH and C_A (Table 3), but they also had substantial dissimilarities in most of the remaining architectural traits, reflected in differences in stemflow yield (Table 3). Syc3 had greater leaf and branch biomass, with more primary and secondary branches, but exhibited less stem lean (5°) than the other individuals (10° and 19° from the vertical axis). The sycamore individual where stemflow was allowed to yield freely into the soil beneath (Syc2) was larger size in terms of DBH , C_A , and Ht , and was more inclined, but had less leaves and less inclined branches compared to Syc1 and Syc3 (Table 3).

Table 3. Aboveground architectural traits and total stemflow yield for the monitoring period (July-October, 2018) for the three sycamore individuals. DBH : diameter at breast height; C_A : projected canopy-crown area; c : canopy cover fraction; Ht : tree height; LAI : leaf area index; S_L : stem lean; $mxBr_a$: maximum branch insertion angle; $avBr_a$: mean branch insertion angle; PBr : number of primary branches per unit area of canopy-crown; SBr : number of secondary branches per unit area of canopy-crown; nL : leaf count per unit area of canopy-crown; L_{BM} : leaf biomass per unit area of canopy-crown; Br_{BM} : branch biomass per unit area of canopy crown. * stemflow yield was predicted in Syc2 as shown in Fig. 3b.

Sycamore individual	DBH (m)	C_A (m ²)	c (%)	Ht (m)	LAI	S_L (°)	$mxBr_a$ (°)	$avBr_a$ (°)	PBr (m ⁻²)	SBr (m ⁻²)	nL (m ⁻²)	L_{BM} (g m ⁻²)	Br_{BM} (g m ⁻²)	Total Stemflow yield (mL
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														m ⁻²)
Syc1	0.37	38.88	93	7.41	5.28	10	55	38	0.08	1.54	499.64	213.60	2557.06	2341.77
Syc2	0.49	53.84	89	11.38	4.53	19	40	28	0.09	1.30	429.53	183.63	2512.81	137.17*
Syc3	0.34	37.34	98	4.06	3.83	5	50	35	0.13	2.11	689.29	294.66	3765.61	3521.61

3.2. Stemflow yield and funnelling ratio

Figure 4. (a) Stemflow yield was significantly higher in sycamore Syc3 ('+' symbol and red line, $y=1.23x-1.45$, $R^2 = 0.69$, $P < 0.01$) compared to Syc1 (empty circles and black line, $y=0.77x-0.88$, $R^2 = 0.63$, $P < 0.01$), with respect to incident rainfall (mm d⁻¹); (b) stemflow funnelling ratio (SFR, unitless) for Syc1 and Syc3. The lower edge of the box corresponds to the 25th percentile data point, while the top edge of the box corresponds to the 75th percentile data point. The line within the box represents the median. The grey area around the box shows the probability density of the data at different values. See online version for colours.

Stemflow yield significantly increased with the incident rainfall in both sycamore individuals (Fig. 4a), it being significantly greater in Syc3 when compared to Syc1. The rainfall threshold to yield stemflow was 1.14 mm d⁻¹ and 1.18 mm d⁻¹ for Syc1 and Syc3, respectively. The total stemflow yield during the monitoring period was 91.05 l and 131.50 l for Syc1 and Syc3, respectively. For both trees and the 35 rainfall events examined, the mean stemflow amount generated per unit projected crown area averaged 1.35 % of the incident rainfall, with a maximum of 3.76 %. Mean stemflow funnelling ratios were above 1.0 in all cases (i.e. more incident rainfall was concentrated around the tree bole than expected had there not been a tree; Fig. 4b) but they did not differ statistically between the two sycamores ($\chi^2=3.46$, $df=1$, $p=0.06$). However, the mean stemflow funnelling ratio was substantially higher in Syc3 (5.16 ± 3.91) than in Syc1 (2.23 ± 2.11).

3.3. Double-funnelling and tree root distribution

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Figure 5. (a) Area within the soil-root zone where subsurface flow had occurred due to double-funnelling and root vertical distribution for sycamores Syc1 (full green circles and dotted line) and Syc2 (full black circles and dotted line). A negative exponential model (lines) was fitted to the measured root area (points) (Gonzalez-Ollauri and Mickovski, 2016). Soil stained with Brilliant Blue FCP dye indicated double-funnelling into soil beneath (b) Syc1 and (c) Syc2. See online version for colours.

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393 In trees that had been sprayed with dye, zones where subsurface flow had occurred due to
394 double-funnelling were successfully identified in the soil close to the stem (Figs. 5b and c).

395 The dye solution mostly infiltrated into the topmost soil layers, reaching maximum depths of
396 350-400 mm b.g.l (Fig. 5a). Specific zones with signs of preferential flow were also
397 identified, and associated with the presence of thicker roots (Figs. 5b-c). The area of soil
398 wetted by double-funnelling (Fig. 5a) was not significantly different between Syc1 and Syc2
399 ($\chi^2=0.18$, $df=1$, $p=0.68$; Fig. 5a). However, we detected a strong positive correlation ($r=0.57$)
400 between the stained area of soil and vertical root distribution, and both decreased with
401 increasing soil depth (Fig. 5a). The root cross-sectional area (Ar) of the two sycamores
402 decreased exponentially with increasing soil depth and had mean rooting depths (i.e. b : soil
403 depth at which 95 % of the roots are located; Gonzalez-Ollauri and Mickovski, 2016; Table
404 1) of 258.9 mm and 275.8 mm for Syc1 and Syc2, respectively.

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407 3.4. Influence of stemflow and its suppression on soil-water dynamics

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411 3.4.1. Soil temperature

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413 Daily soil temperature showed a clear response to stemflow (Fig. 6a) in that it was usually
414 significantly higher ($\chi^2=3143$, $df=1$, $p < 0.01$) under the tree where stemflow was suppressed.

However, under Syc2 with stemflow, a substantial increase in soil temperature was observed following heavy rainfall (i.e. $> 5\text{-}10\text{ mm d}^{-1}$; Fig. 6a) that was not detected in the tree with suppressed stemflow.

3.4.2. Soil moisture

Distinct daily soil moisture (θ_v) peaks were observed under trees with stemflow after heavy rainfall events (i.e. $> 5\text{-}10\text{ mm d}^{-1}$) and following stemflow simulations (Fig. 6b). This response was more pronounced in soil where double-funnelling had occurred, but it was not detected when stemflow was suppressed. Despite this, the θ_v time series did not show significant differences between stemflow and suppressed stemflow ($\chi^2=2.30$, $df=1$, $p=0.13$), not even between the locations with and without signs of double-funnelling ($\chi^2=1.89$, $df=1$, $p=0.17$; Table 2). However, significant differences occurred between soil depths where soil moisture probes were deployed: shallow soil (i.e. 100 mm b.g.l) had significantly higher θ_v compared to deeper soil (i.e. 260 and 400 mm b.g.l; Table 2; $\chi^2=13.09$, $df=2$, $p < 0.01$; Fig. 6b). We excluded soil moisture records from θ_{v-1} (Table 2; Figs. 2c,d), as the moisture probe was dysfunctional (Fig. 6b).

Figure 6. (a) Daily mean soil temperature time series recorded at two points (Table 2) where signs of double-funnelling had occurred under Syc1 (suppressed stemflow) and Syc2 (with stemflow), plotted together with daily mean air temperature and incident daily rainfall data. (b) Daily mean volumetric soil moisture content recorded for Syc1 (suppressed stemflow) and Syc2 (with stemflow), at different areas of the root-soil zone (Table 2), and plotted together with incident rainfall data. (c) Daily mean matric suction recorded at the root-soil zone (Table 2) under Syc1 (suppressed stemflow) and Syc2 (with stemflow). Vertical dot-dash lines indicate stemflow simulation events undertaken after trench excavation (Section 2.5). See online version for colours.

3.4.3. Soil matric suction

Daily soil matric suction (ψ) responded to stemflow markedly over the monitoring period (Fig. 6c) and increased (i.e. became more negative) in both sycamore individuals under dry conditions (i.e. in the absence of rainfall), until it reached the maximum measurable value possible with the tensiometer (-93 kPa; Fig. 6c). However, ψ decreased sharply following heavy rainfall events ($>5 \text{ mm d}^{-1}$) in the sycamore individual that had double-funnelling (Syc2; Fig. 6c). However, Syc1 (i.e. suppressed stemflow) showed no change in ψ (Fig. 6c). The same effect, although of lower magnitude, was observed following stemflow simulations around Syc2 (vertical dot-dash lines in Fig. 6c). As a result, ψ was significantly different between individuals with stemflow, and those where stemflow was suppressed ($\chi^2=44.40$, $df=1$, $p<0.01$). At the end of the observation period, ψ in soil beneath Syc2 decreased towards positive values (i.e. positive pore-water pressure; Fig. 6c). We excluded soil matric suction records after 16/08/2018 (Table 2), as the ψ -2 probe was dysfunctional after this date (Fig. 6c).

3.5. Framework outputs

3.5.1. Effect of tree architectural traits on stemflow yield

Figure 7. (a) Incident rainfall had the highest relative influence (RI) on stemflow yield for Syc1 and Syc3, followed by several architectural traits related to branch dimensions and leaf cover; the white dot within the box represents the median while the grey area around the box shows the probability density of the data at different values (b) Histogram showing the frequency of coefficients of determination (R^2) for the 100 boosted regression trees fitted between aboveground traits against stemflow yield for sycamores Syc1 and Syc3 (c) Regression tree dendrogram for the best performing BRT model fitted to predict stemflow yield from tree architectural traits and incident rainfall. Each tree leaf (i.e. box) indicates the mean response (i.e. stemflow yield in ml), number, and percentage of observations. The darker the colour shade in the tree leaf, the higher is the mean response.

The relative influence (*RI*) of architectural traits (Fig. 7a) on stemflow yield was significantly different ($\chi^2=1225$, $df=15$, $p<0.01$), implying that the measured architectural traits contributed differently to the partition of rainfall into stemflow aboveground. Incident rainfall, which was included as covariate in the BRTs (Section 2.6.1; Fig. 7c), was the most important predictor (*Pg*; $31.48\pm16.02\%$; Fig. 7a). Stem lean (S_L), crown shape (C_{SH}), and biomass (C_{BM}), branch biomass (Br_{BM}) and canopy cover fraction (c) did not influence the production of stemflow yield (Fig. 7a). In the light of the best performing BRT (Fig. 7c) and partial dependence plots (PDPs; supplementary material Figs. S2a-p), incident rainfall and the number of primary branches were strongly and positively correlated with stemflow yield (Figs. 7c, S2a,f) while a strong, negative correlation was observed between stemflow yield and *DBH* (Figs. 7c, S2b). The PDP between stemflow yield and incident rainfall (Fig. S2a) indicated that there was a rainfall threshold of 3.5 mm d^{-1} for the production of stemflow, in contrast with the thresholds of 1.14 and 1.18 mm d^{-1} observed in Fig. 4a (Section 3.2). Beyond rainfall of 3.5 mm d^{-1} , stemflow yield was the same. The remaining aboveground traits did not show clear correlations with stemflow yield (Figs. 7c, S2), in spite of the observed *RI* (Fig. 7a). The fitted BRTs presented a maximum $R^2\pm\text{RMSE}$ of 0.94 ± 19140 , and a mean $\pm$ SD and mode R^2 of 0.19 ± 0.26 and 0.25 , respectively (Fig. 7b).

3.5.2. Effects of stemflow yield, incident rainfall, and air temperature on soil-water dynamics

3.5.2.1. Soil temperature

The BRTs fitted to soil temperature (Eq. 6 - Table 1; Fig. 3) had a high goodness of fit overall (supplementary material Figs. S3a-b; Table 4). The R^2 density function for *ST* (Figs. S3a-b) exhibited negative skewness and a mean value above 0.5 (Table 4). Model quality was significantly higher ($\chi^2=37.94$, $df=1$, $p<0.01$) under Syc1 with suppressed stemflow (Fig. S2a) compared to Syc2 (with stemflow, Fig. S3b). The assessment of the variables' *RI* for all the BRTs fitted to soil temperature (Fig. 8a-b) suggested that air temperature was the most important covariate for predicting soil temperature (Fig. 6a), which was significantly more important than rainfall and stemflow for Syc2 ($\chi^2=92.55$, $df=2$, $p<0.01$). However, air temperature and rainfall were equally important for predicting *ST* when stemflow was

suppressed ($\chi^2=3.67$, $df=1$, $p=0.05$). The influence of incident rainfall and air temperature on soil temperature was corroborated in the partial dependence plots for Syc2 (supplementary material Figs. S4a-c), whereas the effect of stemflow was unclear (Fig. S4b). For Syc1, however, the effect of rainfall on soil temperature was uncertain (Fig. S4a), while air temperature had a more constant influence on soil temperature than that observed beneath Syc2 (Fig. S4c).

Figure 8. Relative influence (RI) of stemflow yield, incident rainfall and air temperature on soil-water dynamics for sycamores Syc1 (suppressed stemflow) and Syc2 (with stemflow) (a-b) soil temperature (c-d) soil moisture (e-f) matric suction. The lower edge of the box corresponds to the 25th percentile data point, while the top edge of the box corresponds to the 75th percentile data point. The line within the box represents the median. The grey area around the box shows the probability density of the data at different values.

3.5.2.2. Soil moisture

The BRTs fitted to soil moisture (θ_v ; Eq. 7 - Table 1; Fig. 3) had a satisfactory goodness of fit in almost all cases under Syc1, where stemflow was suppressed (supplementary material Figs. S3c-h; Table 4). The coefficient of determination (R^2) was also significantly higher ($\chi^2=194.31$, $df=1$, $p<0.01$) under Syc1 (Figs. S3c-d) compared to Syc2 (Figs. S3e-h). The assessment of the variables' RI for all the BRTs fitted to θ_v (Figs. 8c-d) suggested that air temperature was the most important predictor for soil moisture (Fig. 6b), even more so than rainfall and stemflow (Syc1: $\chi^2=205.9$, $df=1$, $p<0.01$; Syc2: $\chi^2=851.6$, $df=2$, $p<0.01$). However and, on the basis of the PDPs (supplementary material Fig. S4), only air temperatures ranging between 10 and 11° C seemed to produce a minor, yet consistent, response on soil moisture (Fig. S4f). The PDPs also showed that the nature of the relationship between rainfall and air temperature with θ_v was unclear for both sycamore individuals (Figs. S4a-f). The same issue was observed between stemflow yield and soil moisture (Fig. S4e).

3.5.2.3. Soil matric suction

The BRTs fitted to soil matric suction (ψ ; Eq. 8 - Table 1; Fig. 3) exhibited, in general, a poor goodness of fit (supplementary material Figs. S3i-j; Table 4). No significant differences occurred between the models fitted under Syc1 (suppressed stemflow) and Syc2 (with stemflow, $\chi^2=0.14$, $df=1$, $p=0.71$). The assessment of the variables' RI for all the BRTs fitted to ψ (Figs. 7e-f) suggested that air temperature was the most important predictor (Fig. 8c), and was significantly more important than rainfall and stemflow (Syc1: $\chi^2=10.54$, $df=1$, $p<0.01$; Syc2: $\chi^2=167.2$, $df=2$, $p<0.01$). On the basis of the PDPs (supplementary material Figs. S4j-l), the influence of rainfall and air temperature on matric suction appeared constant for Syc1. However, these variables had a cyclical effect on matric suction for Syc2 (Figs. S4j-l). The PDPs for stemflow yield showed a negative relationship with matric suction in Syc2, i.e. higher stemflow yields led to lower matric suction (Fig. S4k).

Table 4. Summary from the cross-validation process (Fig. 2) for the 100 boosted regression trees fitted between meteorological variables and soil-water dynamics (Table 1) for the sycamores Syc1 (suppressed stemflow) and Syc2 (with stemflow). *ST*: soil temperature; θ_v : soil moisture; ψ : matric suction; R^2 : coefficient of determination; RMSE: root mean square error; SD: standard deviation.

Variable	Tree	Probe	R^2			Skewness	Kurtosis	SD
			max±RMSE [†]	mean	mode			
Soil temperature	Syc1	ST1	0.95±0.50	0.71	0.79	-1.34	4.96	0.19
	Syc2	ST2	0.94±2.24	0.49	0.71	-0.27	2.08	0.26
Soil moisture	Syc1	θ_v - 2	0.94±0.00	0.39	0.38	0.66	-0.17	2.04
		θ_v - 3	0.84±0.01	0.35	0.38	0.38	-0.12	1.72
	Syc2	θ_v - 4	0.77±0.02	0.18	0.13	0.55	0.52	2.11
		θ_v - 5	0.80±0.02	0.19	0.18	-0.14	0.38	2.38
		θ_v - 6	0.83±0.02	0.17	0.11	-0.14	0.65	2.29
		θ_v - 7	0.60±0.03	0.04	-0.02	-0.14	1.06	3.32
Matric suction	Syc1	ψ - 1	0.99±16.92	0.04	-0.22	1.60	5.69	0.30
	Syc2	ψ - 2	0.95±72.71	0.08	-0.33	0.91	2.55	038

4. Discussion

4.1. Stemflow funnelling above ground

Through our novel framework, we showed that relationships existed between stemflow yield and aboveground tree architecture (Figs. 7a-c). In particular, we demonstrated that a thin trunk and small crown increased stemflow yield and funnelling. Our results also showed that stemflow yield was related to the geometry of the tree's crown, in agreement with Levia and Frost, (2003), Levia et al., (2015) and Yuan et al., (2017). More numerous and steeply angled branches, together with a larger surface area, also increased stemflow production, as found by Levia et al., (2015). However, whilst stemflow yield has been found to be negatively correlated with leaf number (Levia and Frost, 2003; Levia et al., 2015), we showed that more leaves actually increased stemflow (Table 3). This result, which relied on a very small sample size, suggests that leaves could deflect part of the intercepted rainfall towards the woody parts of the canopy, thus contributing to stemflow yield (e.g. Martinez-Meza and Whitford, 1996; Deguchi et al., 2006; Liang et al., 2009; Yuan et al., 2017).

Stemflow yield in sycamores fell within commonly reported values for other woody species across biomes (e.g. Carlyle-Moses et al., 2018) and incident rainfall was the most influential

variable affecting stemflow yield (Yuan et al., 2017). We detected a minimum rainfall threshold of 1.14 mm d⁻¹ required to trigger stemflow (Fig. 4a), and a maximum threshold of 3.5 mm d⁻¹, beyond which stemflow yield was constant (supplementary material Fig. S3a).

Nevertheless, the number of individuals investigated in our study was low, and a wider diversity of tree sizes and architectural traits (e.g. leaf shape and angle of insertion on host branch or bark topography) need examining, so that the model fits between stemflow yield and aboveground architectural traits can be improved. A variety of tree morphologies would also help us to better understand the influence of crown architectural traits on stemflow yield, that we could not disentangle in our study (Figs. 7c, S2). However, our framework was robust enough to account for some of the variability in the linear relationship between stemflow yield and incident rainfall (Figs. 4a, 7a, 7c, S2; Deguchi et al., 2006). Including rainfall as a covariate in the BRTs (Fig. 7c) was useful to gain insights into how rainfall events can be partitioned into stemflow by an array of tree architectural traits at the individual level, and also the hydrological boundaries at which this happens.

The poor BRT fits possibly underline the difficulty of capturing how stemflow is affected by a complex canopy structure (Levia et al., 2015), but a larger tree sample would help to reduce uncertainty. Reliable information about how tree architecture distributes precipitation within the crown to produce stemflow will be especially useful for urban foresters who need to manage stormwater flow around trees that require regular pruning. The type of pruning performed could actually alter the quantity of rainfall that reaches the soil, as well as its transfer belowground (del Campo et al., 2014).

4.2. Double-funnelling

An effective concentration of incident rainfall occurred around the tree bole and in the uppermost soil layers, as the stemflow funnelling ratio was > 1 (Carlyle-Moses et al., 2018). Subsurface flow occurred mainly as matrix flow (Schwärzel et al., 2012; Spencer and van Meerveld, 2016), with some preferential flow observed along coarse woody roots. The dense root system in the topsoil, that was comprised chiefly of thin roots (i.e. diameter < 3 mm), was significantly and positively correlated to double-funnelling. Coarse woody roots visible on the soil surface next to the tree bole may have acted as small dams, causing stemflow to

pond locally and facilitating its infiltration as matrix flow (Mein and Larson, 1973). Ponding could have also been fostered by the hydrophobicity of soil organic matter (Spencer and van Meerveld, 2016), or a higher proportion of silt at the soil surface (unpublished data; Lu and Likos, 2004). The low soil moisture content that we observed, likely reduced the extent of stemflow-derived surface runoff (Liu et al., 2019), that was only noted next to the tree stems (Fig. 2b). Although some preferential flow was observed (Figs. 5b-c), it would be useful to test whether double-funnelling changes from matrix to preferential flow or to surface runoff under different soil hydrological regimes and under different stemflow rates.

Double-funnelling had a clear impact on soil temperature and moisture (Figs. 6a-b) with both variables increasing rapidly after heavy rainfall events (i.e. $> 5 \text{ mm day}^{-1}$ in Syc2). The arrival of water to specific patches of soil, together with peaks in soil temperature could be due to enhanced matrix and preferential flow, and the subsequent increase in microbial activity and respiration (McClain et al., 2003; Kuryakov and Blagodatskaya, 2015) related to the transport of water and nutrients from the canopy to the soil through stemflow (e.g. Germer et al., 2012). Still, further work is required to quantify the effects of double-funnelling on soil respiration fluxes along with the activity of soil microbial communities (e.g. Rosier et al., 2016).

Soil matric suction was significantly modified in response to double-funnelling (Fig. 6c). Between July and August, we observed two clearly defined wetting fronts that only occurred in the soil-root zone under Syc2, suggesting that stemflow can lead to soil matric suction depletion (Liang et al., 2011). In addition, the positive pore-water pressures that developed under Syc2 after the second recorded wetting front were indicative of the formation of a perched water table at the location where the tensiometer was installed (Germer, 2013). The decrease in soil matric suction that we observed can drastically reduce the mechanical strength of plant-soil composite materials (Vanapalli et al., 1996; Gonzalez-Ollauri and Mickovski, 2017b), thus diminishing the mechanical reinforcement provided by the root system in vegetated slopes (Gonzalez-Ollauri and Mickovski, 2016, 2017b, 2017c). It is not yet known to what extent double-funnelling can alter soil hydrological regimes so that soil slippage and landslides could occur. To prevent such potential occurrences, it is necessary to make a choice on planting tree species based not only on mechanical and hydrological traits, but also taking into account aerial architecture and its potential impact on stemflow.

4.3. Effects of stemflow yield, incident rainfall and air temperature on soil-water dynamics

Our framework was useful for detecting effects of stemflow yield and meteorological variables on soil-water dynamics. Air temperature was shown to be the most influential meteorological variable (Fig. 8) and its strong positive correlation with soil temperature could have obscured the effects of rainfall and stemflow yield as predictors of soil temperature. During the summer months, the temperature-dependent, atmospheric demand for water acted as the driver regulating the soil water balance (e.g. Allen et al., 1998; Novick et al., 2016) and hence the dynamics of soil moisture and matric suction in the soil-root zone. However, the effect of rainfall on soil temperature recorded for Syc2, suggested that double-funnelling quickly brought rainfall into the root-soil matrix, and warmed the soil by triggering biogeochemical reactions (Wang et al., 2015; Lloyd and Taylor, 1994; Schindlbacher et al., 2011). In the tree with suppressed stemflow, the effect of rainfall on soil-water dynamics could be related to subsurface flow originated beyond the studied plot, based on the analysis of time lags between rainfall and soil-water dynamics (e.g. supplementary material – Fig. S5; Bogaard, 2001; Bestland et al., 2009).

Surprisingly, stemflow yield was the least important covariate influencing soil-water dynamics in Syc2. The strong correlation between rainfall and stemflow (Figs. 4a, 7a, S2a) may have obscured the relationship with stemflow in the BRTs (model covariates should be independent from each other, Bruce and Bruce, 2017). However, in our study, there were limitations to experimental design because stemflow yield was not directly measured for Syc2, but was predicted for Syc2 using BRTs fitted to a small dataset. Furthermore, Syc2 had substantial architectural differences with respect to the individuals used to fit stemflow yield BRT models (Table 3), which likely led to the poor BRT fits (Fig. S3). Therefore, to clarify the effect of stemflow on soil-water dynamics, it is essential to quantify stemflow yield for a larger sample and longer periods, using flow meters or tipping bucket gauges (e.g. Levia et al., 2010; Spencer and van Meerveld, 2016; del Campo et al., 2018) before allowing stemflow to funnel belowground.

Some limitations occurred when using BRTs, in particular, when we evaluated the relationships between predictors and response variables in the partial dependence plots (PDPs; supplementary material Figs. S2, S4). The discretisation of the response variables by

BRTs only enabled us to observe weak predictor-response interactions in the PDPs, as opposed to the array of effects discussed for the time series records (Section 4.2; Fig. 6) as well as for the relationships between architectural traits and stemflow yield (Section 4.1). To circumvent this issue, we encourage the incorporation of alternative statistical models able to generate continuous outputs (e.g. random forest; [Breiman, 2001](#)) in future versions of our framework.

5. Conclusion

We demonstrated how a novel numerical framework and experimental approach can be used to examine the effect of tree aboveground architecture on stemflow yield and its influence on soil-water dynamics. In the light of our observations and findings, it can be concluded that:

- The number of branches, their insertion angle, leaf number, and stem basal diameter influenced stemflow yield within specific rainfall thresholds.
- Funnelling of stemflow beneath the soil surface occurred as matrix flow and was significantly and positively correlated with the vertical root distribution.
- Soil-water dynamics were distinctly different with and without stemflow.
- Soil matric suction was negatively affected by stemflow yield, but air temperature was the most influential covariate affecting soil-water dynamics likely due to its strong correlation to evapotranspiration during the summer season.
- The discretisation of the response variables by boosted regression trees only enabled us to observe weak predictor-response interactions, as opposed to the array of effects observed in this study.

In spite of the study limitations discussed above, such as small sample size and differences between individuals, the proposed framework and experimental approach provide a good basis for future research contributing to our knowledge of how stemflow generated aboveground triggers major responses in soil-water dynamics belowground.

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6. References

- Alexander, H., Arthur, M., 2010. Implications of a predicted shift from upland oaks to red maple on forest hydrology and nutrient availability. *Can. J. For. Res.* 40(4), 716-726
- Allen, R. Pereira, L., Raes, D., Smith, M., 1998. Crop evapotranspiration guidelines for computing crop water requirements. *FAO Irrigation and drainage paper* No 56.
- Barthélémy, D., Caraglio, Y., 2017. Plant Architecture: A Dynamic, Multilevel and Comprehensive Approach to Plant Form, Structure and Ontogeny. *Annals of Botany*, 99, 375-407.
- Bestland, E. Milgate1, S., Chittleborough, D., VanLeeuwen, J., Pichler, M., Soloninka, L., 2009. The significance and lag-time of deep through flow: an example from a small, ephemeral catchment with contrasting soil types in the Adelaide Hills, South Australia. *Hydrol. Earth Syst. Sci.*, 13, 1201–1214
- Bialkowski, R., Buttle, J. M., 2015. Stemflow and throughfall contributions to soil water recharge under trees with differing branch architectures. *Hydrol. Process.*, 29, 4069-4082.
- Blozan, W., 2008. Tree measuring guidelines of the eastern native tree society. *Bull. East. Native Tree Soc.* 1 (1), 3–10.
- Böhm, W., 1979. *Methods of Studying Root Systems*. Springer Verlag, Berlin.
- Bogaard, T., 2001. Analysis of hydrological processes in unstable clayey slopes. *Nederlandse, Geografische Studies*
- Breiman L, Freidman J, Olshen R, Stone C (1984) *Classification and regression trees*. Wadsworth, Belmont, CA.
- Breiman, L., 2001. Random forests. *Mach. Learn.* 45, 5–32.
- Bruce, P., Bruce, A., 2017. *Practical statistics for data scientists*. O’Reilly, Sebastopol, CA.

742 Carlyle-Moses, D. E., Iida, S., Germer, S., Llorens, P., Michalzik, B., Nanko, K., Tischler, A.,
 743 Levina, D. F., 2018. Expressing stemflow commensurate with its ecohydrological importance.
 744 *Advances in Water Resources*, 121, 472-479.
 745 Carlyle-Moses, D. E., Price, A. G., 2006. Growing-season stemflow production within a
 746 deciduous forest of southern Ontario. *Hydrol. Process.*, 20, 3651-3663.
 747 Côté, J. F., Fournier, R. A., Egli, R., 2011. An architectural model of trees to estimate forest
 748 structural attributes using terrestrial LiDAR. *Environmental Modelling and Software*, 26,
 749 761-777.
 750 Deguchi, A., Hattori, S., Park, H., 2006. The influence of seasonal changes in canopy
 751 structure on interception loss: application of the revised Gash model. *J. Hydrol.* 318, 80–102.
 752 del Campo, A. D., Fernandes, T. J. G., Molina, A. J., 2014. Hydrology-oriented (adaptive)
 753 silviculture in semiarid pine plantation: How much can be modified the water cycle through
 754 forest management? *Eur. J. Forest. Res.*, 133(5), 879-984.
 755 del Campo, A. D., Gonzalez-Sanchis, M., Lidon, A., Ceacero, C. J., Garcia-Prats, A., 2018.
 756 Rainfall partitioning after thinning in two low-biomass semiarid forests: Impact of
 757 meteorological variables and forest structure on the effectiveness of water-oriented
 758 treatments. *J. Hydrol.*, 565, 74-86.
 759 Efron, B., 1979. Bootstrap methods: another look at the Jackknife. *Ann. Stat.* 1, 1–26.
 760 Feng, H., Liu, Y., 2015. Combined effects of precipitation and air temperature on soil
 761 moisture in different land covers in a humid basin. *J. Hydrol.*, 531(3), 1129-1140.
 762 Germer, S., 2013. Development of near-surface perched water tables during natural and
 763 artificial stemflow generation by babassu palms. *J. Hydrol.*, 507, 262-272.
 764 Germer, S., Werther, L., Elsenbeer, H., 2010. Have we underestimated stemflow? Lessons
 765 from an open tropical forest. *J. Hydrol.*, 395(3-4), 169-179.
 766 Germer, S., Zimmermann, A., Neill, C., Krusche, A. V., Elsenbeer, H., 2012.
 767 Disproportionate single-species contribution to canopy-soil nutrient flux in an Amazonian
 768 rainforest. *For. Ecol. Manage.*, 267, 40-49.
 769 Gonzalez-Ollauri, A., Mickovski, S.B., 2016. Using the root spread information of pioneer
 770 plants to quantify their mitigation potential against shallow landslides and erosion. *Ecol. Eng.*
 771 95, 302–315.
 772 Gonzalez-Ollauri, A., Mickovski, S.B., 2017a. Hydrological effect of vegetation against
 773 rainfall-induced landslides. *J. Hydrol.* 549, 374–387.
 774 Gonzalez-Ollauri, A., Mickovski, S.B., 2017b. Plant-soil reinforcement response under
 775 different soil hydrological regimes. *Geoderma* 285, 141–150.

776 Gonzalez-Ollauri, A., Mickovski, S.B., 2017c. Plant-Best: A novel plant selection tool for
 777 slope protection. *Ecol. Eng.*, 106, 154-173.

778 Gonzalez-Ollauri, A., Mickovski, S.B., 2017d. Shallow landslides as drivers for slope
 779 ecosystem evolution and biophysical diversity. *Landslides*, 14(5), 1699-1714.

780 Greenwell, B. M., 2017. pdp: An R Package for Constructing Partial Dependence Plots. *The*
 781 *R Journal*, 9(1), 421--436. URL [https://journal.r-project.org/archive/2017/RJ-2017-](https://journal.r-project.org/archive/2017/RJ-2017-016/index.html)
 782 [016/index.html](https://journal.r-project.org/archive/2017/RJ-2017-016/index.html).

783 Hamdan, K., Schmidt, M. G., 2012. The influence of bigleaf maple on chemical properties of
 784 throughfall, stemflow, and forest floor in coniferous forest in the Pacific Northwest. *Can. J.*
 785 *For. Res.*, 42(5), 868-878.

786 Head, K.H., Epps, R.J., 2011. *Manual of Soil Laboratory Testing: Permeability. Shear*
 787 *Strenght and Compressibility Tests*, vol. 2. CRC Press, Boca Raton.

788 Herwitz, S. R., 1986. Infiltration-excess caused by stemflow in cyclone prone tropical
 789 rainforest. *Earth Surface Processes and Landforms*, 11, 401-412.

790 in Forested Ecosystems. In: *Forest Hydrology and Biogeochemistry: Synthesis of Past*
 791 *Research and Future Directions, Ecol. Stud. Ser.* Vol. 216, edited by Levias, D. F., Carlyle-
 792 Moses, D. E., Tanaka, T., pp. 483-498, Springer, Heidelberg.

793 Johnson, M. S., Jost, G., 2011. Ecohydrology and Biogeochemistry of the Rhizosphere
 794 Johnson, M. S., Lehmann, J., Riha, S., Krusche, A. V., Richey, J. E., Ometto, J.P., Couto, E.
 795 G., 2008. CO₂ efflux from Amazonian headwater streams represents a significant fate for
 796 deep soil respiration. *Geophysical Research Letters*, 35, L17401.

797 Johnson, M. S., Lehmann, J., 2006. Double-funnelling of trees: Stemflow and root-induced
 798 preferential flow. *Ecoscience*, 13(3), 324-333.

799 Kattge, J. et al., 2011. TRY – a global database of plant traits. *Global Change Biology*, 17,
 800 2905-2935.

801 Kuhn, M. et al., 2018. caret: Classification and Regression Training. R package version
 802 6.0.81. <https://CRAN.R-project.org/package=caret>

803 Kuzyakov, Y., Blagodatskaya, E., 2015. Microbial hotspots and hot moments in soil: Concept
 804 & review. *Soil Biology & Biochemistry*, 83, 184-199.

805 Lemon, P., E., 1956. A Spherical Densimeter For Estimating Forest Overstory Density.
 806 *Forest Sci.* 2, 314-320.

807 Levias, D. F., Frost, E.E., 2003. A review and evaluation of stemflow literature in the
 808 hydrologic and biogeochemical cycles of forested and agricultural ecosystems. *J. Hydrol.*,
 809 274(1-4), 1-29.

810 Levia, D. F., Germer, S., 2015. A review of stemflow generation dynamics and stemflow-
811 environment interactions in forests and shrublands. *Rev. Geophys.*, 53, 673-714.

812 Levia, D. F., Herwitz, S. R., 2000. Physical properties of stemflow water in relation to
813 leachate dynamics: Implications for nutrient cycling. *Can. J. For. Res.*, 30(4), 662-666.

814 Levia, D. F., Keim, R. F., Carlyle-Moses, D. E., Frost, E. E., 2011. Throughfall and stemflow
815 in wooded ecosystems. In: *Forest Hydrology and Biogeochemistry: Synthesis of Past
816 Research and Future Directions, Ecol. Stud. Ser.* Vol. 216, edited by Levia, D. F., Carlyle-
817 Moses, D. E., Tanaka, T., pp. 425-443, Springer, Heidelberg.

818 Levia, D. F., van Stan, J. T., Inamdar, S. P., Jarvis, M. T., Mitchell, M. J., Mage, S. M.,
819 Scheick, C. E., McHale, P. J., 2012. Stemflow and dissolved organic carbon cycling:
820 Temporal variability in concentration, flux, and UV-vis spectral metrics in a temperate
821 broadleaved deciduous forest in the eastern United States. *Can. J. For. Res.*, 42(1), 207-216.

822 Levia, D. F., Van Stan, J. T., Mage, S. M., Kelley-Hauske., 2010. Temporal variability of
823 stemflow volume in a beech-yellow poplar forest in relation to tree species and size. *J.*
824 *Hydrol.*, 380(1-2), 112-120.

825 Levia, D. F., Michakzik, B., Bischoff, S., Richter, S., Legates, D. R., 2015. Differential
826 stemflow yield from European beech saplings: The role of individual canopy structure
827 metrics. *Hydrol. Process.*, 29(1), 43-51.

828 Liang, W. L., Kosugi, K., and Mizuyama, T., 2009. A three-dimensional model of the effect
829 of stemflow on soil water dynamics around a tree on a hillslope, *J. Hydrol.*, 366(1-4), 62-75.

830 Liang, W. L., Kosugi, K., and Mizuyama, T., 2015. Soil water redistribution process around a
831 tree on a hillslope: The effect of stemflow on the drying process, *Ecohydrology*, 8, 1381-
832 1395.

833 Liang, W. L., Kosugi, K., Mizuyama, T., 2007. Heterogenous soil water dynamics around a
834 tree growing on a steep hillslope. *Vadose Zone J.*, 6(4), 879-889.

835 Liang, W., Kosugi, K., Mizuyama, T., 2011. Soil water dynamics around a tree on a hillslope
836 with or without rainwater supplied by stemflow. *Water Resour. Res.* 47 (W02541).

837 Liu, J., Engel, B. A., Wang, Y., Wu, Y., Zhang, Z, Zhang, M., 2019. Runoff response to soil
838 moisture and micro-topographic structure on the plot scale. *Scientific Reports*, 9(1), 2532.

839 Lloyd, J., Taylor, J. A., 1994. On the temperature dependence of soil respiration. *Functional*
840 *Ecology*, 8, 315-323.

841 Lu, N., Godt, J., 2013. *Hillslope Hydrology and Stability*. Cambridge University Press, New
842 York.

843 Lu, N., Likos, W.J., 2004. *Unsaturated Soil Mechanics*. John Wiley and Sons, Hoboken, US.

844 Martinez-Meza, E., Whitford, W. G., 1996. Stemflow, throughfall and channelization of
845 stemflow by roots in three Chihuahuan desert shrubs. *J. Arid Environ.*, 32(3), 271-287.

846 McClain, M. E. Boyer, E. W., Dent, C. L., Gergel, S. E., Grimm, N. B., Groffman, P. M.,
847 Hart, S. C., Harvey, J. W., Johnston, C. A., Mayorga, E., 2003. Biogeochemical hot spots and
848 hot moments at the interface of terrestrial and aquatic ecosystems. *Ecosystems*, 6(4), 301-
849 312.

850 Mein, R.G., Larson, C.L., 1973. Modeling infiltration during steady rain. *Water Resour. Res.*
851 9 (2), 384–394.

852 Meinzer, F. C., Lachenbruch, B., Dawson, T., E. (Eds.), 2011. *Size- and Age-Related*
853 *Changes in Tree Structure and Function*. Tree Physiology, Vol. 4. Springer, Heidelberg.

854 Miyata, S. K., Kosugi, K., Gomi, T., Mizuyama, T., 2009. Effects of forest floor coverage on
855 overland flow and soil erosion on hillslopes in Japanese cypress plantation forests. *Water*
856 *Resour. Res.*, 45, W06402.

857 Molina, A., del Campo, A. D., 2012. The effects of experimental thinning on throughfall and
858 stemflow: a contribution towards hydrology-oriented silviculture in Aleppo pine plantations.
859 *For. Ecol. Manage.*, 269, 2016-213.

860 Nespoulus, J., Merino-Martin, L., Monier, Y., Bouchet, D., Ramel, M., Dombey, R.,
861 Viennois, G., Mao, Z., Zhang, J., Cao, K., Le Bissonnais, Y., Sidle, R., Stokes, A., 2019.
862 Tropical forest structure and understorey determine subsurface flow through biopores formed
863 by plant roots. *Catena*, In Press.

864 Norris, J.S. et al., 2008. *Slope Stability and Erosion Control: Ecotechnological Solutions*.
865 Springer, Doordrecht.

866 Novick, K. A. et al., 2016. The increasing importance of atmospheric demand for ecosystem
867 water and carbon fluxes. *Nature Climate Change*, NCLIMATE3114

868 Partington, D., Brunner, P., Frei, S., Simmons, C. T., Werner, A. D., Therrien, R., Maier, H.
869 R., Dandy, G. C., Fleckenstein, J. H., 2013. Interpreting streamflow generation mechanisms
870 from integrated surface-subsurface flow models of a riparian wetland and catchment. *Water*
871 *Res. Research*, 49, 5501-5519.

872 Pasta, S., de Rigo, D., Caudullo, G., 2016. *Acer pseudoplatanus* in Europe: distribution,
873 habitat, usage and threats. In: San-Miguel- Ayanz, J., de Rigo, D., Caudullo, G., Houston
874 Durrant, T., Mauri, A. (Eds.), *European Atlas of Forest Tree Species*. Publ. Off. EU,
875 Luxembourg.

876 Preti, F., Dani, A., Laio, F., 2010. Root profile assessment by means of hydrological,
 877 pedological and above-ground vegetation information for bio-engineering purposes. *Ecol.*
 878 *Eng.* 36, 305–316.

879 R Core Team, 2018. R: A language and environment for statistical computing. R Foundation
 880 for Statistical Computing, Vienna, Austria. URL <https://www.R-project.org/>.

881 Rosier, C. L., Levia, D. F., van Stan, J. T., Aufdenkampe, A., Kan, J., 2016. Seasonal
 882 dynamics of the soil microbial community structure within the proximal area of tree boles:
 883 Possible influence on stemflow. *European Journal of Soil Biology*, 73, 108-118.

884 Rossi, R. E., Mulla, D. J., Journel, A. G., Franz, E. H., 1992. Geostatistical Tools for
 885 Modeling and Interpreting Ecological Spatial Dependence. *Ecological Monographs*, 62(2),
 886 277-314.

887 Schindlbacher, A., Rodler, A., Kuffner, M., Kitzler, B., Sessitsch, A., & Zechmeister-
 888 Boltenstern, S. (2011). Experimental warming effects on the microbial community of a
 889 temperate mountain forest soil. *Soil biology & biochemistry*, 43(7), 1417–1425.
 890 doi:10.1016/j.soilbio.2011.03.005

891 Schwärzel, K., Ebermann, S., Schalling, N., 2012. Evidence of double-funnelling effect of
 892 beech by visualization of flow pathways using dye tracer. *J. Hydrol.* 470-471, 184-192.

893 Shumway, R. H., Stoffer, D. S., 2017. *Time Series Analysis and Its Applications: With R*
 894 *Examples*. Springer Texts in Statistics. Springer, Heidelberg.

895 Spencer, S. A., van Meerveld, H. J., 2016. Double funnelling in a mature coastal British
 896 Columbia forest: spatial patterns of stemflow after infiltration. *Hydrol. Process.*, 30, 4185-
 897 4201.

898 Tanaka, N., Levia, D., Igarashi, Y., Yoshifuji, N., Tanaka, K., Tantasirin, C., Nanko, K.,
 899 Suzuki, M., Kumagai, T., 2017. What factors are most influential in governing stemflow
 900 production from plantation-grown teak trees? *J. Hydrol.*, 544, 10-20.

901 Tardio, G., Gonzalez-Ollauri, A., Mickovski, S.B., 2016. A non-invasive root distribution
 902 analysis methodology from a slope stability approach. *Ecol. Eng.* 97, 46–57.

903 Therneau, T., Atkinson, B., 2018. rpart: Recursive Partitioning and Regression Trees. R
 904 package version 4.1-13. <https://CRAN.R-project.org/package=rpart>
 905 US.

906 van Stan, J. T., Levia, D. F., Inamdar, S. P., Lepori-Buib, M., Mitchell, M. J., 2012. The
 907 effects of phenoseason and storm characteristics on throughfall solute washoff and leaching
 908 dynamics from a temperate deciduous forest canopy. *Science of the Total Environment*, 430,
 909 48-58.

Vanapalli, S.K., Fredlund, D.G., Pufahl, D.E., Clifton, A.W., 1996. Model for the prediction of shear strength with respect to soil suction. *Can. Geotech. J.* 33, 379–392.

voor de Poorte, P., 2018. Retrieved 01 04, 2019, from PEDROX, live weather from Catterline: <<http://www.pedrox.com>>.

Wang, L., Manzoni, S., Ravi, S., Rivero-Iregui, D., Caylor, K., 2015. Dynamic interactions of ecohydrological and biogeochemical processes in water-limited systems. *Ecosphere*, 6(8), Article 133

Wang, W. et al., 2016. Evaluation of air–soil temperature relationships simulated by land surface models during winter across the permafrost region. *The Cryosphere*, 10, 1721-1737.

Whitford, W. G., Anderson, J., Rice, P. M., 1997. Stemflow contribution to the ‘fertile island’ effect in creosotebush, *Larrea tridentate*. *Journal of Arid Environments*, 35, 451-457.

Wolf, D., Carson, E.A., Brown, R.H., 1972. Leaf area index and specific leaf area determinations. *J. Agron. Educ.* 1, 24–27.

Wright, I. J., Reich, P. B., Cornelissen, J. H. C., Falster, D. S., Groom, P. K., Hikosaka, K., Lee, W., Lusk, C. H., Niinemets, U., Oleksyn, J., Osada, N., Poorter., H., Warton, D. I., Westoby, M., 2005. Modulation of leaf economic traits and trait relationships by climate. *Global Ecol. Biogeogr.*, 14, 411–421.

Wua, L., McGechanb, M.B., McRobertsc, N., Baddeleya, J.A., Watsona, C.A., 2007. SPACSYS: Integration of a 3D root architecture component to carbon, nitrogen and water cycling—Model description. *Ecological Modelling*, 200, 343-359.

Yuan, C., Gao, G., Fu, B., 2017. Comparisons of stemflow and its bio-/abiotic influential factors between to xerophytic shrub species. *Hydrol. Earth Syst. Sci.*, 21, 1421-1438.

Zimmermann, A., Zimmermann, B., 2014. Requirements for throughfall monitoring: the roles of temporal scale and canopy complexity. *Agric. For. Meteorol.* 189–190, 125–139.